

1 **Using morphospaces to understand tafoni development**

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10

11 **Abstract**

12

13 Tafoni research has tended to focus on issues around definition and differences rather
14 than trying to develop general concepts for understanding the nature of tafoni. This
15 paper uses the concepts of fitness landscapes and morphospaces to develop a
16 standardized and dimensionless phase space within which to represent, visualize and
17 analyze a dataset of 800 tafoni collected from Antarctica. Within this phase space it
18 is possible to identify clustering of tafoni forms and to illustrate how tafoni
19 development is constrained by a relational hierarchy of rock structure, processes and
20 geometry or form.

21

22 **Keywords:** tafoni, development, fitness landscapes, morphospaces

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24

25 **1. Introduction**

26

27 Tafoni have been the source of debate in geomorphology since the first identification
28 and proposed explanation of this distinctive form (see Groom et al., in Press).
29 Unfortunately, key issues arise again and again in the literature as the supposed
30 ‘distinctiveness’ of this form eludes definition. This elusiveness means that any
31 definitive statement on the characteristics and diagnostic processes of this form are
32 almost impossible to delineate. Specifically, the debate hovers around issues of scale
33 (are ‘small’ tafoni the same as ‘large’ tafoni?), development (do small tafoni become
34 large tafoni and is there a distinct developmental sequence to tafoni formation, do
35 they represent self-organization?), and process-form relationships (is there a
36 diagnostic set of processes that cause tafoni to develop and maintain the form?).
37 Research tends to focus on either one or all of these issues. The underlying
38 assumption of form indicating process and changes in form indicating changes in
39 process is at the heart of the measurement and analysis of tafoni.

40

41 *1.1 What are tafoni?*

42

43 There appears to be a number of terms relating to ‘hollows’ developed in bedrock, the
44 most common of which (in English) are ‘honeycomb’ or ‘alveolar’ weathering and
45 ‘tafoni’ (e.g. Evelpidou et al., 2010); the whole often being referred to as ‘cavernous
46 weathering’ (e.g. Turkington and Phillips, 2004; Viles, 2005). The terms ‘honeycomb
47 weathering’ and ‘cavernous weathering’ seem to be the catch-all terms for the
48 creation of “small caves” (Evelpidou, et al., 2010) or “caverns” (Turkington and
49 Phillips, 2004) developed by differential weathering in rock. In many of these studies
50 the distinction between form terminology appears to be almost solely related to size

51 rather than to actual form or process (Groom et al., in Press). This, thus, leaves the
52 question as to whether alveolar weathering is but a precursor of tafoni and/or whether
53 the size distinction is simply a product of the host lithology. According to Evelpidou
54 et al (2010, p. 34), following Penck (1894), “honeycomb weathering formations
55 bigger than 0.5 m are defined as Tafoni, whereas formations smaller than 0.5 m are
56 defined as Alveoles”; seemingly the whole defined as ‘honeycomb weathering’.
57 Mustoe (1982) provides extensive information regarding nomenclature and some of
58 the confusion resulting from non-standardization of terminology. Cavernous
59 weathering is often used to encompass all the other terms (e.g. French and
60 Guglielmin, (2000) refer to tafoni as an attribute of cavernous weathering) but may
61 also be considered as an entity in its own right (e.g. Dragovich, 1967). Thus, the
62 question arises as to quite what **are** tafoni and where, if at all, do they fit within the
63 spectrum of other associated terms?

64

65 To some extent, many of the background components of this discussion have been
66 covered by Viles (2005) and the reader is directed towards this excellent review. Key
67 within the study of Viles (2005, p. 1471) is the opening statement: “Understanding the
68 initiation, development and significance of landforms remains a central issue in
69 geomorphology.” Indeed, the whole issue regarding initiation of these weathering
70 forms remains an enigma (Boxermann, 2005, p.79). However, to the above points
71 must also be added the caveat that ‘terminology’ (see Hall et al., 2012) requires we all
72 understand the same thing through the use of specific terms; this does not appear to be
73 the case with respect to the terms used here. In part, this may well underpin the
74 observation by Turkington (2004, p.128) that “as more information has been
75 presented their (tafoni and alveoli) possible origins, rather than being clarified, seem

76 to have become more confused.” Perhaps some of this confusion is related to our use
77 of terms and that perhaps the forms these terms refer to are either a continuum (rather
78 than discrete) or **are** discrete and not part of a continuum (see Inkpen, 2005, for a
79 discussion on these issues within geomorphology).

80

81 Viles (2005) clearly uses the term ‘cavernous weathering’ to encompass a number of
82 forms (notably tafoni and alveoli – see her Fig. 1) as too do Turkington and Phillips
83 (2004). Here it is argued, much as discussed elsewhere for other processes (see Hall et
84 al., 2012), that the foundational terminology ‘cavernous weathering’ itself creates
85 confusion – is it (cavernous **weathering**) the ‘process’ (as actually implied by the
86 term) or the product (the ‘cavern’) and if it is the ‘cavern’ then quite what does this
87 encompass; or is it implying (as does appear to be the case) both process **and** form?
88 Where, as it would appear here, both process and form are included within the term,
89 then this creates many issues (much as it has in nivation – see Thorn and Hall, 2002)
90 as to the conflating of process and form within one term. Thus, while Viles (2005)
91 makes an excellent case for the advances made regarding ‘cavernous weathering’,
92 notably the self-organizational attributes of form development, the very real problems
93 of both terminology and process remain. Indeed, Viles (2005, p. 1472) alludes to this
94 very issue where it is stated that the overall outcome “rather than providing a
95 consensus viewpoint or indicating a clearly developing research field, seems to be
96 ‘mine are different to yours’.” This may, though, be either the very issue or that
97 various workers, simply because the terminology is failing us, do not recognize that
98 they are indeed dealing with comparable forms.

99

100 1.2 *Form and process relationships*

101

102

103 There clearly is much confusion regarding the nature of the formative weathering (or,
104 rather ‘rock decay’: see Hall, et al., 2012) – essentially everything from chemical to
105 physical to physico-chemical processes, and almost any combination thereof. This, in
106 itself, need not be a problem as this paper argues. Indeed, the very extent and variety
107 of suggested processes is not necessarily unexpected given that cavernous weathering
108 is azonal in occurrence (Turkington and Phillips, 2004) and found in a variety of
109 lithologies (see Mustoe, 1982, Table 1). Given the variety of identified causative
110 processes, the product appears to be a classic ‘convergence of form’, as already noted
111 by Turkington and Phillips (2004, p. 666). That being the case, then perhaps the
112 question is one of why do these different processes produce the same end result?

113

114 In turn this may beg the question, as to whether the processes are any different in their
115 **effect** on the rock; the effect is to solely disassociate the constituent materials. The
116 **nature** of that disassociation may well be controlled more by lithology than process.
117 In other words, if ‘flaking’ (the effect) is the outcome, it can be the product of a
118 variety of causes (wetting-drying, thermal stresses, salt weathering, freeze-thaw,
119 chemical processes, etc) acting alone or in combinations. If that were the case then it
120 may be less important as to what the formative process was and, in turn, suggests rock
121 properties may play the key role (see Hall et al., 2012). It may also be, however, that
122 it is the relations between the form and process and the factors that control these
123 relations, rather than the dominance or otherwise of any particular component, that is
124 the essential aspect to understanding any generalized conceptualisation of tafoni
125 evolution.

126

127 Burridge and Inkpen (2015) highlight this in the mathematical model of tafoni
128 development. In this paper rock properties provide the context within which process
129 operate to produce the tafoni form. One might argue that, given convergence of form
130 resulting from a multitude of identified processes, then maybe the focus of research
131 should be on underlying factors such as rock properties that can constraint
132 development or, in a more subtle conceptual framework, the relations between factors
133 that may be canalizing development.

134

135 This paper suggests that this seemingly unsatisfactory state of affairs may help in
136 developing a novel conceptual framework within which to interpret tafoni. This paper
137 suggests that viewing tafoni within the conceptual framework of fitness landscapes
138 and morphospaces permits ‘fuzziness’ in definitions within the context of the factors
139 that constrain development and which define the parameter phase spaces for tafoni
140 development. In order to advance this argument we first outline the nature of fitness
141 landscapes and morphospaces. Secondly, we identify the three key factors and their
142 parameter phase spaces that constrain tafoni as derived from the existing literature.
143 We highlight the importance of a relational view of these factors for defining the
144 canalizing outcome in phase space. By canalizing we mean that the parameter spaces
145 confine and guide the development of forms along specific pathways. As individual
146 tafoni become increasingly embedded within these developmental pathways, the
147 constraints imposed by these parameter spaces become increasingly difficult to
148 overcome. Lastly, using this conceptual framework we illustrate how it might be used
149 to interpret simple dimensional measurement of tafoni derived from Dronning Maud
150 Land in the Antarctic. From this analysis we are able to show that tafoni inhabit a

151 restricted area of the phase space and that the detailed analysis of dimensions within
152 this zone may not yield any additional information about process and form
153 relationships. If appropriate then this conceptual framework suggests which aspects of
154 form-process relationships should be the focus of further research into tafoni
155 development.

156

157 **2. Fitness Landscapes and Morphospaces**

158

159 Within the biological literature, as noted by McGhee (2007), the concept of ‘adaptive
160 landscapes’ originates with Wright (1932) who used the concept to visualize the
161 fitness of genes, although he coined the term ‘fitness landscape’ for his visualization
162 (Figure 1). The adaptive landscape represents all the possible combinations of genes
163 that an organism might produce. From these possible combinations, those that
164 actually existed could be identified and plotted. The fittest of the existing
165 combinations could be thought of as peaks rising from the relatively unfit surface. In
166 Figure 1, for example, there are two possible ‘fit’ peaks and Wright proposed that
167 evolution by natural selection would force gene combinations to climb the nearest
168 peak, always moving gene combination towards fitter variants. Movement is also
169 informed by local conditions, so even if a nearby peak is lower than the lowest peak
170 globally, variants will move towards that nearest, lower peak. The topography of a
171 fitness landscape provides a roadmap of possible evolutionary pathways. Adaptive
172 landscapes have also been defined in hyperdimensions by Kaufmann (1995);
173 Gaverilets and Gravener (1997) and Gavrilets (2003); and with the latter suggesting
174 that the complex and multiple nature of parameters affecting adaption result in a
175 relatively flat but multidimensional landscape covered with holes. The holes represent

176 locations where planes of fitness intersect and so are regions or clusters of hyperspace
177 where fit gene combinations can occur. In the above author's landscape, evolution can
178 be 'smooth' within the clusters but 'jumpy' as gene combinations move from one
179 cluster to another through 'extradimensional bypasses' (Gavrilets (1997, p.311).
180 Within geomorphology Phillips (2009) outlined a similar vision of landscape
181 evolution with his concept of Landscape Evolution Space (LES), an n-dimensional
182 space or hypervolume representing the resources, energy, and the parameters
183 available for landscape development. Conceptually, any landscape should be capable
184 of being located within this hypervolume and its trajectory or development mapped
185 out in the same space. Inkpen and Petley (2007) offer something similar in their
186 analysis of landform development.

187

188 Theoretical morphospaces are not the same as adaptive landscapes but are related to
189 them (McGhee, 2007). Developed by Raup (1966, 1967), a morphospace can be
190 described as a hyperspace of geometries, with axes representing different
191 morphological traits, that represent all the forms possible if these traits are
192 systematically altered. Within a morphospace the axes represent dimensions and
193 form; the resulting surface is a representation of how frequently that form appears.
194 The morphospace provides an indication of the forms that occur in reality, and
195 importantly, those that do not. The two types of space can be linked if the distribution
196 of forms in the morphospace have adaptive significance. Raup (1966), for example,
197 studied the form of ammonoids and identified that there was a distinct pattern to this
198 distribution in morphospace. Chamberlain (1976, 1981) through experimental work
199 on models, found that the two regions of morphospace created by ammonoid forms
200 were those where swimming efficiency was maximised. Form was linked to

201 adaptation. Regions of morphospace do not necessarily match to peaks that optimise
202 a specific function, but rather, as in the research into plant morphospaces by Niklas
203 (1997, 2004), the peaks represent geometries that minimize several functional
204 problems. This highlights that fitness is always a concept that needs to be thought of
205 in multidimensional terms.

206

207 Combining the two spaces, McGhee (2007) develops an argument that they can be
208 used to explore the constraints that exist upon development. Fig. 2 illustrates the
209 concept that development is constrained by a series of factors; geometric, functional,
210 phylogenetic and developmental in the case of organisms. McGhee defined the
211 geometric and functional constraints as extrinsic, being imposed by the laws of
212 physics and chemistry, whilst phylogenetic and developmental constraints were
213 intrinsic, imposed by the biology of specific organisms. Assuming a form can be
214 defined by a set of measurements then the total possible set of forms will be defined
215 by points in a hyperspace as in Fig. 2. Within this set of possible forms will be a
216 subset of forms that represent all geometrically possible forms (GPF in Fig. 2). Co-
217 ordinates outside of this region of hyperspace represent geometrically impossible
218 forms. McGhee defines the boundary between these regions of hyperspace as the
219 ‘geometric constraint boundary’. Nested within the GPF are two other regions,
220 functionally possible forms (FPF) and functionally impossible forms (FIP). The result
221 is a clearly defined subset of hyperspace that demarcates the region of possible forms
222 given the nested series of constraints. Importantly, the extrinsic constraints remain
223 constant and define rigid boundary conditions, whilst the intrinsic constraints vary
224 with taxon and so are more flexible in the boundaries they prescribe. Recent work on
225 the simulation of vegetated aeolian landscapes (Baas, 2007; Baas and Nield, 2007,

226 2010; Nield and Baas, 2008) provide illustrations of the clustering of forms in a
227 simulated parameter space.

228

229 Brierley (2010), building upon Brierley and Fryirs (2005), identifies a structuring of
230 explanations concerning landscape development in a similar manner: identifying
231 geologic, climatic and anthropogenic memory. Brierley views these three types of
232 memory as imposing differing limits upon landscape development. Geologic memory
233 imposes boundary conditions within which contemporary landscape-forming
234 processes continue to operate, whilst climatic memory controls the nature and
235 effectiveness of geomorphic processes. Anthropogenic memory alters the fluxes of
236 sediment and flows in the landscape. Brierley (2010) is at pains to point out that these
237 factors operate collectively and variably across different time frames despite the
238 temptation to view them as hierarchical. This suggests that explanation in
239 geomorphology is structured around sets of parameters that continually constrain the
240 possible forms and the potential pathways of their development.

241

242 There may be a basis for seeing a conceptual analogy between morphospaces and
243 fitness spaces and the concept of strange attractors (Phillips, 1999, 2003). Both sets of
244 concepts discuss mapping system properties in a phase space within which certain
245 portion of space are more likely to be populated than others. Within Phillips'
246 discussion, strange attractors are areas of phase space to which evolutionary
247 trajectories are drawn. In the language of morphospaces this means that the zone of
248 the attractor will define a region of particular form characteristics. Within this region
249 there will be a highly proportion or percentage of measured individuals. The attractor
250 need not represent an evolutionary basin but rather the range of forms that can be

251 taken given variations in constraining properties. The most frequent forms represent
252 the most common outcome but other forms nearby could represent the manifestation
253 of slightly different relations between constraining properties and yet still define a
254 basin of attraction.

255

256 Application of both concepts to geomorphology does, however, face key problems
257 that mirror those found within biology. Fitness spaces need to be defined in relation to
258 some concept of fitness that then needs to be translated into empirical, measurable
259 terms for defining the extent of morphospace. Identifying ‘fitness’ implies having a
260 clear concept of the expected trajectory of a form and a clear understanding of the
261 basis for this trajectory. Similarly, the plotting of individuals within a morphospace
262 requires the identification and quantification of important characteristics of form. Our
263 traditional ways of thinking about tafoni from affect what forms we identify in the
264 field and what we deem important to measure. Likewise, technical constraints such as
265 the type of equipment available, its measurement resolution and the ability to
266 consistently measure a highly variable natural phenomenon will all impact upon the
267 nature and quality of data available to characterise forms.

268

269 **3. The Spaces of Tafoni**

270

271 Combining fitness landscapes and morphospaces it is possible to analyse the
272 parameters that define the morphospaces of tafoni and then the manner in which these
273 forms change as tafoni form clusters and developmental sequences. The
274 morphospaces that combine to constrain tafoni formation, development and form are
275 structural, process-based and geometric. These three morphospaces are related to each

276 as in Fig. 3, in a nested hierarchy with each additional space constraining the potential
277 location of tafoni in the morphospaces. It is important to bear in mind that the figure
278 is a representation of multidimensional spaces of rock structure, process and geometry
279 and their relations in two-dimensions; it is a visual aid to interpretation. Fig. 3
280 illustrates that tafoni development is constrained by rock structure but rock structure
281 itself is not sufficient to determine whether tafoni develop. Rock structure instead
282 defines a section of morphospace within which tafoni could develop. Potential tafoni
283 development in this morphospace is further constrained by other factors as discussed
284 below. Collectively these form the hierarchy of constraining factors as illustrated in
285 the figure. Burrige and Inkpen (2015) outline a similar hierarchical structure to
286 modelling tafoni development. Rock properties provide the context within which
287 processes operate to produce a geometry of form which then feedbacks to process and
288 affects rock properties.

289

290 Structural or rock property constraints refer to the various parameters associated with
291 rock properties that have been identified in the past as being associated with tafoni
292 formation. These include inherent weaknesses in the rock, fractures, cracks, as well as
293 porosity, permeability and rock composition. It is within this structurally defined
294 constrained morphospace that processes of weathering and erosion operate and,
295 importantly, interact with each other and with the structural parameters. Structural
296 morphospace may constrain the potential for tafoni to develop but it is not sufficient
297 on its own to determine whether tafoni will develop.

298

299 Tafoni are inherently about the relations between parameter spaces. For process-
300 defined morphospace it is not the specific processes that are important but rather the

301 nature of the relations between these processes and between these processes and
302 structural parameters. Processes capable of inducing stresses in the near-surface of the
303 rock, which then result in the differentiation of the surface and subsurface properties,
304 are how process-defined morphospace and structural morphospace interact. This
305 means that a range of processes can be vital for tafoni formation. It is not a specific
306 process that causes tafoni to develop but rather it is the result of process relations, in
307 conjunction with rock properties, that produces surface and subsurface differentiation
308 and near-surface stress. Further, this is not a static relationship. Processes and
309 structure interact and in so doing change the nature of that interaction. This means
310 that the morphospaces evolve as well. The structural constraints are initially set very
311 broadly. Adding the process relations produces a refinement of which parts of the
312 spaces are able to develop tafoni. The ongoing interaction between the two further
313 refines this space of potential development and can even expand the spaces of
314 potential as structural properties are altered at the micro-level to become increasingly
315 conducive to tafoni development.

316

317 An outcome of the complicated relationship between structure and process is the
318 development of a distinct geometry to the resultant form. This is the geometric space,
319 a further constraining morphospace. Once the characteristic tafoni form begins to
320 develop there is an interaction with the processes causing surface and subsurface
321 differentiation. The nature of this relationship determines the development of the
322 geometry of the form that in turn affects the dynamics of the structural and process
323 relationships. This further constrains the spaces of tafoni development as well as
324 altering the nature of structural and process spaces to redefine the locations of

325 potential tafoni development. Combined these three spaces produce a nested hierarchy
326 of potential spaces for tafoni.

327

328 Conceptually, the interaction of the three spaces creates broad regions or clusters
329 where tafoni could develop. These clusters need not be contiguous. This means that
330 tafoni of different sizes and shapes are all tafoni formed through the relations between
331 these three parameter spaces, just formed at different intersections of these parameter
332 spaces. This also means that there is not necessarily a developmental sequence from
333 small through middle-sized to large tafoni. The size distribution need not represent a
334 developmental sequence but rather a different combination of relations.

335

336 This means that it could be that different studies have revealed different clusters of
337 tafoni and so different locations of potentiality in the relations between these
338 parameter spaces. Once tafoni are initiated then they will develop into forms
339 constrained by the morphospace. Although the potential forms may be varied there
340 will be limits, boundaries, to these forms. It may be that small tafoni will always
341 remain small as their development is confined within a specific region of tafoni
342 morphospace. Small tafoni can not suddenly jump across morphospace and explore
343 the region inhabited by large tafoni. Likewise large tafoni may initially develop
344 rapidly as the relations between structure, process and geometry permit the rapid
345 exploration of potentiality in that region of morphospace. Once trapped along a
346 particular developmental pathway, however, it may be that the rate of growth slows as
347 the limits to that particular pathway are reached. Wright described this process as
348 channelization of forms, as outlined in section 2. A deep cavern, for example, may be
349 too deep for differentiation between surfaces to occur as weathering products can not

350 be removed to permit further erosion. Conceptually this is limiting the space of
351 potential development for a tafoni as it develops and alters the relations between the
352 morphospaces of the three parameters.

353

354 **4. Illustration of Interpretation of Tafoni Space**

355 Tafoni were measured in Dronning Maud Land, Antarctica in the Austral summers of
356 2008/08 and 2010/11. Tafoni were measured on nunataks on the Ahlmannryggen
357 (Ahlmann Ridge), specifically on nunataks, Vesleskarvet (Northern Buttress; 71°40'S,
358 2°51'W), Lorenzenpiggen (71°45'S, 2°50'W), Grunehogna (72°02'S, 2°48'W),
359 Flarjuven Bluff (72°01'S, 3°24'W) and Robertskollen (71°27'S, 3°15'W). The rock in
360 the area is Precambrian in origin and the exposures measured were of the
361 Borgmassivet Intrusives comprising doleritic and dioritic sills. Measurements were
362 made on 40 rock faces, starting sampling at the central point of each rock face and
363 then measuring the dimensions of the tafoni away from the centre of the rock face
364 until 10 tafoni had been measured. Dimensions were measured using a set of callipers
365 and undertaken by the same observer to ensure consistency in the field definition of
366 length, width and depth.

367

368 The tafoni dataset was converted to dimensionless values using width/length and
369 depth/width ratios and a phase space constructed using these dimensionless
370 parameters as axes. The data were converted to dimensionless values to analysis form
371 changes within the phase space rather than focusing on changes in the size of the
372 tafoni. If the form of the tafoni, i.e. the relative dimensions of length, width and depth,
373 did not change as it grew then more and more tafoni would occupy the same area of
374 phase space. Fig. 4 illustrates the distribution of tafoni in this dimensionless phase

375 space with cells along the x and y axes of 0.2 units. The contours represent the
376 percentage of the tafoni in the dataset of 800 individuals occupying specific areas of
377 the phase space. Although these dimensionless ratios have been used in tafoni
378 research before they have not been used to map the distribution of tafoni in such as
379 phase space. Fig. 4 shows a single peak to the distribution of tafoni at around 0.6-0.8
380 units of both the width/length and depth/width axes with a relatively smooth and
381 continuous decrease around this peak in the occurrence of tafoni. There seems to be a
382 tail in the distribution in the direction of higher depth/width ratios suggesting that
383 there are a number of tafoni becoming deepening whilst retaining a form consistent
384 with those tafoni in the peak area. The single peak and the relatively even spread of
385 tafoni away from it might imply the peak represents the end point of an evolutionary
386 or developmental sequence for tafoni. The relative frequencies of tafoni in the phase
387 space would, if this were the case, represent the stages in tafoni evolution with the
388 peak being the most frequent and final stage. Tafoni not in the peak might represent
389 individual tafoni that had not yet developed to their final form or tafoni where the
390 relations between rock structure, process and geometry in this environment were not
391 as fully expressed as in the peak.

392

393 The contiguous nature and relatively plateau-like nature of the frequency surface in
394 phase space may, however, suggest that the differences between a tafoni at the peak
395 and one in the sub-peak areas may not mean that different processes are at work or
396 that tafoni are at different stages of development. Rather the different zones represent
397 the differing expression of the same set of rock structure, process and geometry
398 relationships. This means that differences in form can not necessarily be interpreted to
399 mean differences in how the tafoni form, only differences in the relative importance

400 of each factor. This means that the exact position of an individual tafoni in the phase
401 space need not reflect major changes in constraining factors nor in the relations that
402 produce its final form. A cloud of individual tafoni positions may reflect the same
403 relations just expressed slightly and insignificantly differently.

404

405 Analysis of variance suggests that there is no statistically significant difference
406 between the tafoni in the three cells forming the plateau region in terms of length and
407 width but depth increases significantly between these cells (analysis of variance with
408 an F value of 0.8 for length and 0.69 for width, both statistically not significant and an
409 F value of 9.43 for depth statistically significant at $\alpha=0.01$). Moving away from the
410 plateau there are statistically significant differences in length, width and depth as
411 might be expected as the ratios change. The nature of the change is consistent in that
412 length, width and depth increase alone in specific areas of the morphospace so
413 significant changes result from the increase in size of a single dimension rather than
414 decreases in size.

415

416 Fig. 5 illustrates the variation in statistically significant increases in length, width and
417 depth of tafoni in the morphospace. Tafoni in this dataset have a limited depth of 10-
418 15 cm whatever the form of the tafoni. This implies that there is a vertical limit to
419 tafoni development, suggesting that the relations between the three key factors only
420 operate within a 10-1cm depth from the rock surface. This depth-limit to relationships
421 between factors was also found in the mathematical modelling of tafoni in Burridge
422 and Inkpen (2015) and may suggest that tafoni development is a depth-limited
423 process. The bottom left of the phase space is dominated by tafoni with high average
424 lengths compared to tafoni in every other part of the phase space. This suggests that

425 tafoni in this section of the phase space are elongated and may represent either
426 controls on shape through rock structure or the coalescence of tafoni lengthwise. The
427 right-hand side of the phase space at mid D/W ratios has high average widths for
428 tafoni compared to tafoni across the rest of the phase space. This suggests that tafoni
429 move to this portion of phase space through widening rather than overall growth in
430 dimensions, as the average tafoni length is not significantly different from the tafoni
431 in the rest of the phase space.

432

433 The morphospace produced illustrates the constraining nature of the three parameters.
434 Rock properties form the common context within which the tafoni develop and the
435 single peak in morphospace implies that this constraint usually produces a single,
436 characteristic set of forms. Process is constrained to the process specific to this
437 environment and the geometry of form seems to be highly constrained to a limited set
438 of ratios between the dimensions measured. The morphospace produced represents
439 the range of tafoni form produced within this rock type, in this weathering
440 environment and provides a template against which to map other tafoni from other
441 environments as well as tafoni of larger and smaller dimensions. If the tafoni from
442 other environments map into a similar zone then this implies that the relations
443 between rock properties, process and geometry are consistent across environments
444 and so forms converge into a specific region of morphospace. If, however, tafoni
445 from other environments map to a different region of morphospace then this implies
446 that there are significant differences in how the relationships between the factors are
447 expressed in different environments. In this case there is a basis for claiming some
448 tafoni are different from others and to question the common terminology to describe
449 them.

450

451

452 **5. Conclusion**

453

454 Expressing the data within a standardized and dimensionless phase space allows the
455 researcher to visualize the forms within a common setting. This can help the
456 researcher to identify patches of the phase space where forms cluster, and to provide a
457 definition of the characteristics that define these clusters in terms of dimensional
458 relationships. Comparing the location of clusters between studies could help to
459 identify if there is a common pattern to clustering within this phase space or if the
460 location and nature of the clusters vary with each study. This will help to distinguish
461 and define tafoni that present common patterns in form relationships and tafoni whose
462 form relationships express their site-specific nature. Making this distinction will help
463 researchers define forms which could be classified as tafoni in any environment as
464 opposed to forms that exhibit tafoni-like tendencies but which cluster in a different
465 part of the phase space. ‘Mine is different from yours’ becomes less of a problem as
466 here is a way of visualizing if and by how much mine is different from yours and if
467 the difference might be significant.

468

469 The role of form and process, as well as the relative importance of other factors, can
470 be analysed using the hierarchical model of morphospaces presented above. The
471 central importance of rock structure defines the limits to the range of forms possible
472 and so could be viewed as the overarching control on the potential for tafoni to
473 develop. Whether tafoni develop or not is not solely determined by rock structure
474 however. The relationships between rock structure and weathering and erosion are

475 vital for determining if tafoni develop and which areas of the phase space the forms
476 inhabit. Producing stress in the near-surface is the key outcome that affects tafoni
477 production and the evolving relationships between processes and form, tightly
478 constrained by structure, establishes the developmental and geometric relationships
479 that are expressed by the forms measured. This could mean that different processes
480 produce different clusters in phase space and so process identification may be aided
481 by mapping these clusters. It may be, however, that the clusters are so broad, as in this
482 example, that the differentiation between processes is not feasible. This could imply
483 that the range of forms produced within the constraints of the morphospace defined by
484 rock structure is potentially large as subtle variations in process-form relationships
485 can be expressed by a wide range of dimensional outcomes. With only one set of data
486 it is difficult to assess if this is a general characteristic of tafoni but setting the
487 discussion within this common framework would enable these key hypotheses to be
488 tested.

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490

491 **6. Acknowledgements**

492

493 The authors would like to thank Ian Meiklejohn, Christel Hansen, Michael Loubser
494 and Werner Nel for the tafoni datasets that were from the ‘Landscape Processes in
495 Antarctic ecosystems’ project funded by the South African NRF. The authors would
496 also like to thank two anonymous reviewers for their constructive comments.

497

498 **7. References**

499

500 Baas, A.C.W., 2007. Complex systems in Aeolian geomorphology. *Geomorphology*,
501 91, 311-331.
502

503 Baas, A.C.W., Nield, J.M., 2007. Modelling vegetated dune landscapes. *Geophysical*
504 *Research Letters* 34, p.06405.
505

506 Baas, A.C.W., Nield, J.M., 2010. Ecogeomorphic state variables and phase-space
507 construction for quantifying the evolution of vegetated Aeolian landscapes. *Earth*
508 *Surface Processes and Landforms* 35, 717-731.
509

510 Boxerman, J.Z., 2005. The evolutionary cycle of the tafoni weathering pattern on
511 sandstone at Bean Hollow Beach, Northern California. *Geological Society of America*
512 *Abstracts with Programs* 37(4), 79.
513

514 Brierley, G.J., 2010. Landscape memory: the imprint of the past on contemporary
515 landscape forms and processes. *Area* 42, 76-85.
516

517 Brierley, G.J., Fryirs, K.A., 2005 *Geomorphology and river management: applications*
518 *of the river styles framework*. Blackwell, Oxford.
519

520 Burrige, J., Inkpen, R., 2015. Formation and arrangement of pits by a corrosive gas.
521 *Physical Review E* 91, 022403.
522

523 Chamberlain, J.A., 1976. Flow patterns and drag coefficient of cephalopod shells.
524 *Palaeontology* 19, 539-563.

525

526 Chamberlain, J.A., 1981 Hydromechanical design of fossil cephalopods in M.R. House
527 and J.R. Senior (eds,) The Ammonoidea. Systematics Association, London. pp.289-
528 336.

529

530 Dragovich, D., 1967. Flaking, a weathering process operating on cavernous rock
531 surfaces. Geological Society of America Bulletin 78, 801-804.

532

533 Evelpidou, N, Leonidopoulou, D. and Vassilopoulos, A., 2010. Tafoni and alveole
534 formation. An example from Naxos and Tinos Islands. In Evelipou, N., de Figueiredo,
535 T., Mauro, F., Tecim, v. and Vassilopoulos, A. (eds,) Natural heritage from East to
536 West. Case studies from 6 EU Countries. Springer. 35-42.

537

538 French, H.M., Guglielmin, M., 2000. Cryogenic weathering of granite northern
539 Victoria land, Antarctica. Permafrost and Periglacial Processes 11, 305–314.

540

541 Gavrilets, S., 1997. Evolution and speciation on holey landscapes. Trends in Ecology
542 and Evolution 12, 307-12.

543

544 Gavrilets, S., 2003. Evolution and speciation in a hyperspace: The roles of neutrality,
545 selection, mutation and random drift. In, Crutchfield, J.P. and Schuster P. (eds,)
546 Evolutionary dynamics. Oxford University Press, Oxford.

547

548 Gavrilets, S., Gravner, J., 1997. Percolation on the fitness hypercube and the
549 evolution of reproductive isolation. Journal of Theoretical Biology 184, 51-64.

550 Glossary of Geology. 2005. Fifth edition. Neuendorf, K.K.E. and Mhel, Jr. ,J.P. (eds)

551 Guglielmin, M., Cannone, N., Strini, A., Lewkowicz, A., 2005. Biotic and abiotic
552 processes on granite weathering landforms in a cryotic environment, Northern
553 Victoria Land, Antarctica. *Permafrost and Periglacial Processes* 16, 69–85.

554 Groom, K., Allen, C.D., Mol, L., Paradise, T.R., Hall, K., In Press. Defining tafoni:
555 Re-examining terminological ambiguity for cavernuous rock decay. *Progress in*
556 *Physical Geography*.

557 Hall, K., Thorn, C., Sumner, P., 2012. On the persistence of ‘weathering’.
558 *Geomorphology* 149, 1-10.

559 Inkpen, R., 2005. *Science, philosophy and physical geography*. Routledge, London.
560

561 Inkpen, R., Petley, D., 2001. Fitness spaces and their potential for visualizing change
562 in the physical landscape. *Area* 33, 242-251.
563

564 Kauffman, S., 1993. *Origins of Order: Self-organization and Selection in Evolution*.
565 Oxford University Press, New York.

566 Kaufmann, S., 1995 *At home in the universe*. Penguin, London.
567

568 Kejonen, A., Kielosto, S., Lahti, S.I., 1988. Cavernous weathering forms in Finland.
569 *Geografiska Annaler* 70A (4), 315–321.
570

571 McGhee, G.R., 2007. The geometry of evolution: Adaptive landscapes and theoretical
572 morphospaces. Cambridge University Press, Cambridge.
573

574 Mustoe, G.E., 1982. The origin of honeycomb weathering. Geological Society of
575 America Bulletin 93, 108–115.
576

577 Neild, J.M., Baas, A.C.W., 2008. The influence of different environmental and
578 climatic conditions on vegetated Aeolian dune landscape development and response.
579 Global and Planetary Change 64, 76-92.
580

581 Niklas, K.J., 1997. The evolutionary biology of plants. University of Chicago Press,
582 Chicago.
583

584 Niklas, K.J., 2004. Computer models of early land plant evolution. Annual Review of
585 Earth and Planetary Sciences 32, 47-66.
586

587 Ollier, C.D., Bourman , R.P., 2002. Flared Slopes, Footholds, and the Retreat of
588 Overhanging Slopes: Examples of Convergent Landform Development. Physical
589 Geography 23, 321-334.
590

591 Owen, A.M., Athersat, W.J., Mylroie, J.E., Mylroie, J.R., 2005. A morphological
592 comparison of tafoni vs. flank margin caves, San Salvador Island, Bahamas.
593 Geological Society of America Abstracts with Programs 38 (7) 62.
594

595 Penck, A. 1894. Morphologie der Erdoberfläche. Stuttgart, Engelhorn, Vol I, 214

596

597 Phillips, J.D., 1999. Earth surface systems: complexity, order and scale. Blackwell,
598 Oxford.

599

600 Phillips, J.D., 2003. Sources of nonlinearity and complexity in geomorphic systems.
601 Progress in Physical Geography 27, 1-23.

602

603 Phillips, J.D., 2009. Landscape evolution space and the relative importance of
604 geomorphic processes and controls. Geomorphology 109, 79-85.

605

606 Raup, D.M., 1966. Geometric analysis of shell coiling: General problems. Journal of
607 Palaeontology 40, 1178-1190.

608

609 Raup, D.M., 1967. Geometric analysis of shell coiling: Coiling in Ammonoids.
610 Journal of Palaeontology 41, 43-65.

611

612 Thorn, C.E., Hall, K., 2002. Nivation and cryoplanation: the case for scrutiny and
613 integration. Progress in Physical Geography 26, 533-550.

614

615 Turkington, A., 2004. Cavernous weathering. In: Goudie, A.S. (ed.), Encyclopaedia of
616 Geomorphology. Routledge, New York. 128–130.

617

618 Turkington, A., Phillips, J.D., 2004. Cavernous weathering, dynamical instability and
619 self-organization. Earth Surface Process and Landforms 29, 665–675.

620

621 Viles, H.A., 2005. Can stone decay be chaotic? In Stone Decay in the Architectural
622 Environment, Turkington AV (ed). Geological Society of America Special Paper 390,
623 11–16.

624

625 Walker, L.N., Mylroie, J.E., Walker, A.D., Mylroie, J.R., 2008. The Caves of Abaco
626 Island, Bahamas: keys to geologic timelines. Journal of Cave and Karst Studies 70,
627 108–119.

628

629 Wright, S., 1932. The role of mutation, inbreeding, crossbreeding, and selection in
630 evolution Proceedings of the Sixth International Congress on Genetics 1, 355–65.

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634 **Figure Captions**

635 Figure 1 Illustration of fitness landscape (Modified from Wright, 1932)

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637 Figure 2 Illustration of concept of development within a morphospace
638 constrained by series of factors (Modified from McGhee, 2007)

639

640 Figure 3 Relationship between structural, process-based and geometric factors
641 in morphospaces as a nested hierarchy

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643 Figure 4 Morphospace of percentage frequency of tafoni occurrence for specific
644 width/length and depth/width ratios

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646 Figure 5 Illustration of changes in tafoni length, width and depth across
647 morphospace

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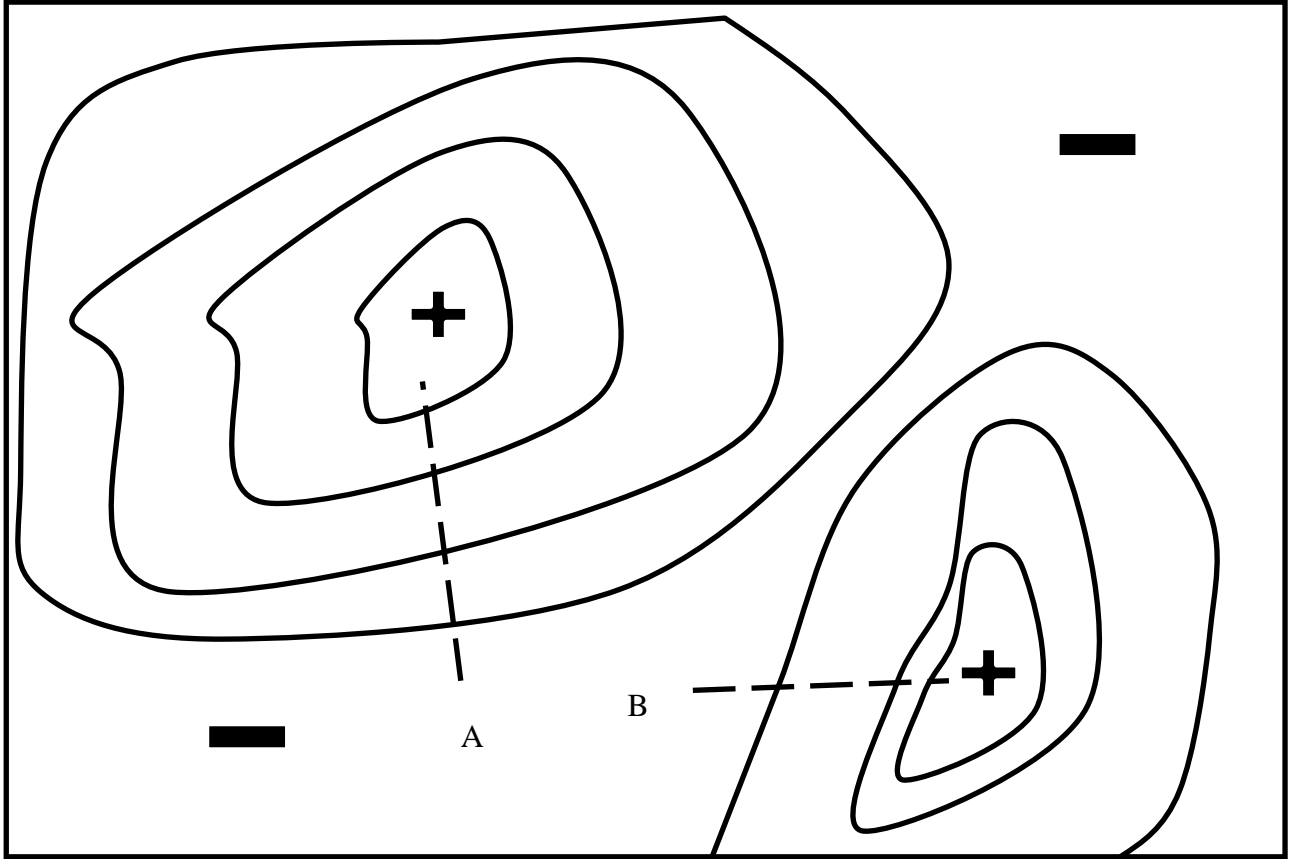
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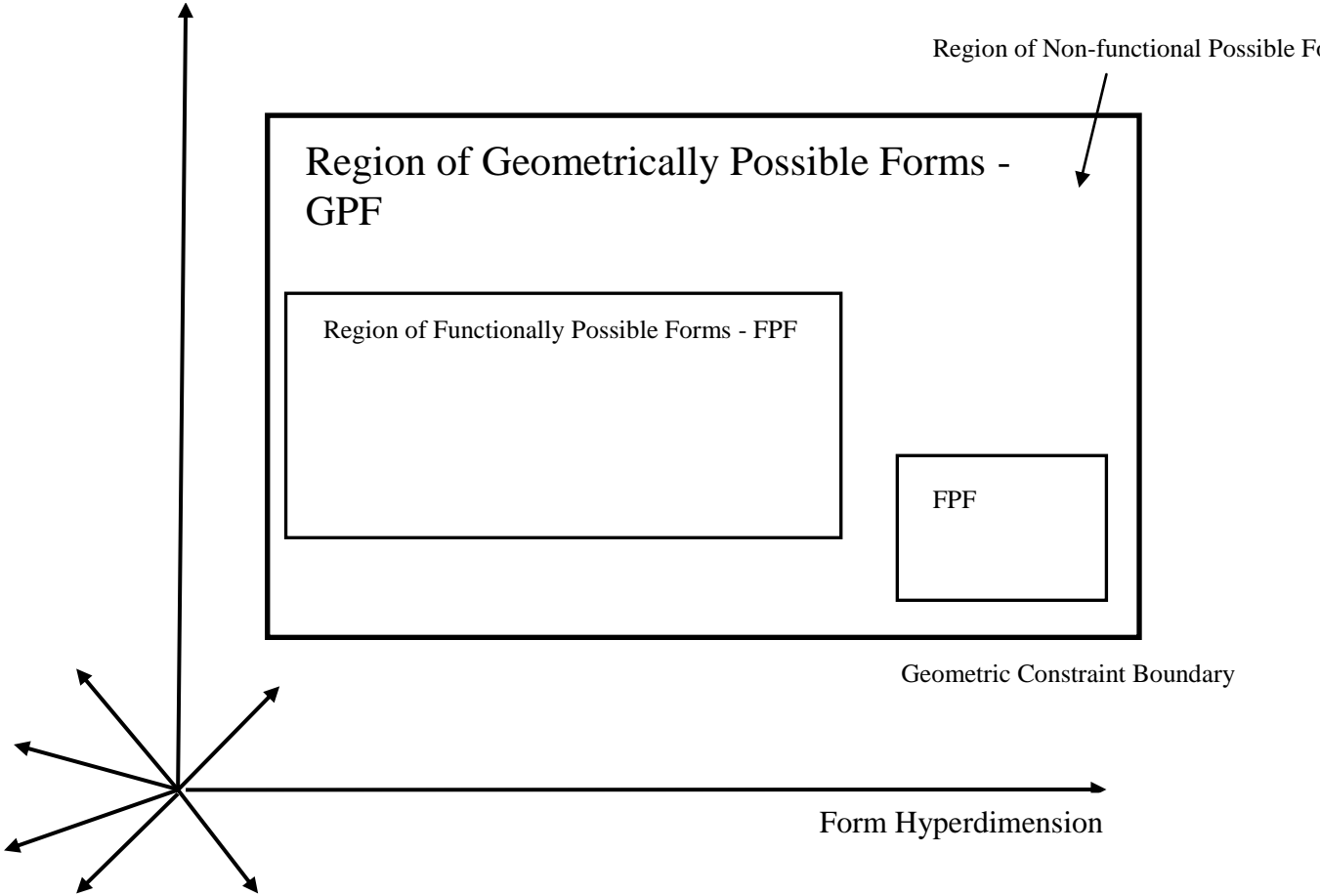
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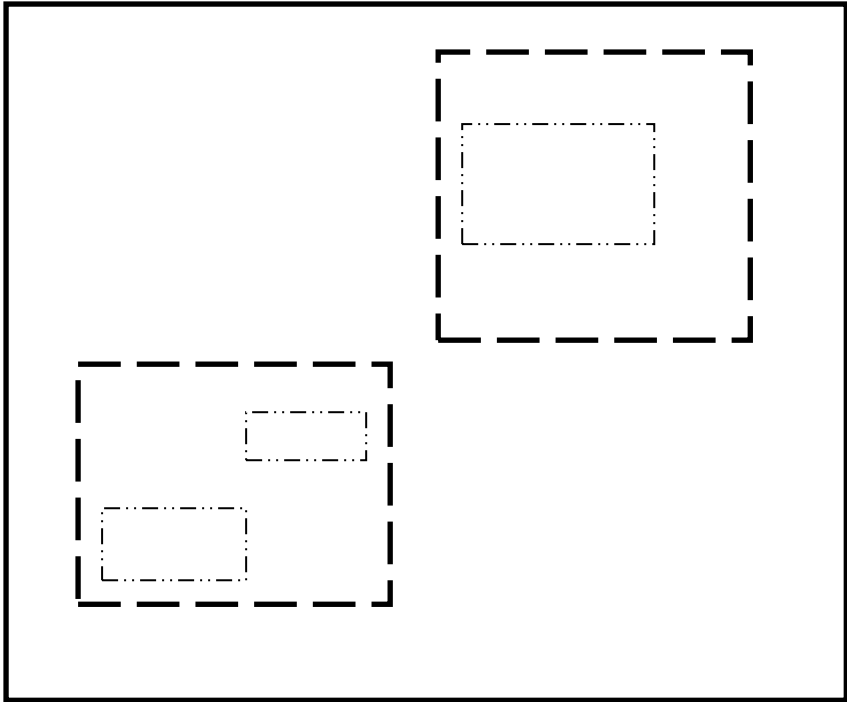


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Total Hyperspace of Form Dimensions (axes represent different hyperdimensions)



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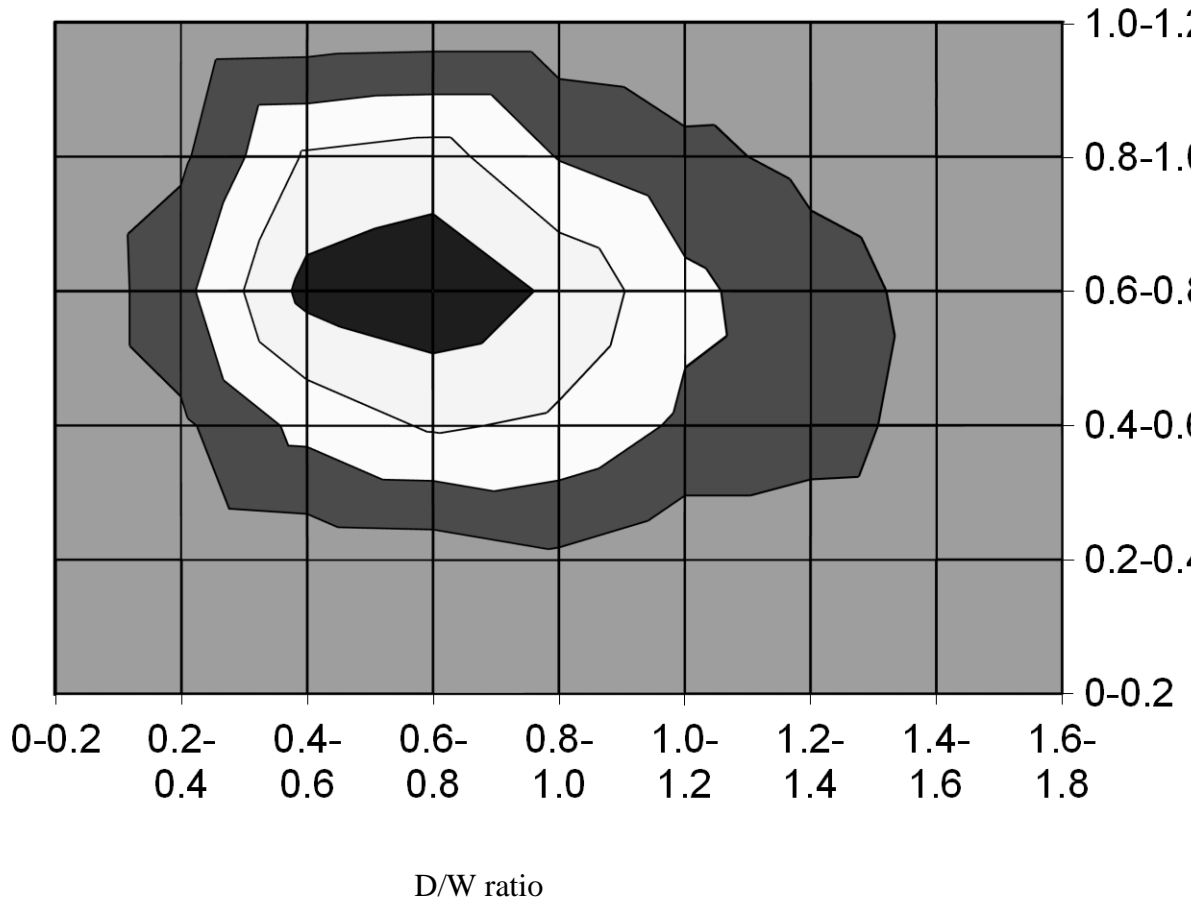
Structurally defined morphospace



Process defined morphospace



Geometrically defined morphospace



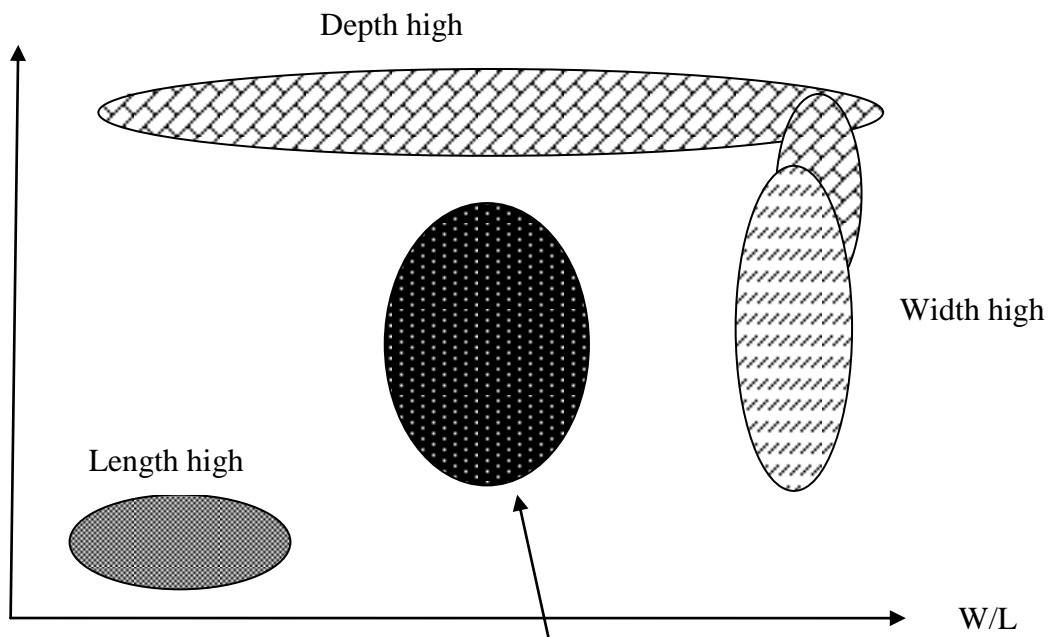
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Plateau - no significant difference in length

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Highlights

- Morphospace of tafoni from Antarctica dataset defined using dimensionless parameters
- Clustering of tafoni identified in morphospace and potential developmental explanation discussed
- Tafoni development constrained by relational hierarchy of rock structure, processes and geometry